SAR Remote Sensing of Snow Parameters in Norwegian Areas — Current Status and Future Perspective

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Abstract—The paper presents results from a series of European and national projects on remote sensing of snow parameters. Currently, satellite borne synthetic aperture radar (SAR) data are only available at C-band frequencies. Other frequencies such as L-band or Ku-band may be favorable in several snow applications, but current C-band SAR may still be used and further developed to a more mature level. In particular, the advent of wide swath SAR data have provided frequent data sets at medium spatial resolution, that can be used to monitor snow parameters operationally.

We will present results from a snow cover area monitoring service developed for Norway and Sweden. The service, which is based on Envisat ASAR wide swath data, produces snow cover maps on average 3-4 times per week. The resulting time series gives a unique data set for studying the snow cover as it rapidly retreats during the melting season, and is of high value to hydro power companies.

Snow water equivalent (SWE) is the key parameter for hydrological applications. Norut IT has developed a technique using repeat pass interferometry to measure SWE, based on the linear relationship between the change in SWE and the change in interferometric phase. The technique has been demonstrated, but scarceness of usable interferometric baseline pairs have so far not allowed wide spread applicability of the technique.

A future SAR using Ku-band frequency as carrier will maybe solve the problem of retrieving SWE. Since backscatter at Ku-band frequency is more sensitive to SWE, it is good hope that robust SAR methods can be invented for this purpose. It will, however, be extremely important for the scientific society to validate the retrieval algorithms against in-situ data. The authors have developed an innovative validation concept using ground-penetrating radars at the same carrier frequencies as the space borne SARs to validate EO data in an efficient manner. The concept has been studied at C-band frequencies on glaciers at Svalbard, and we hope to build a similar platform for Ku-band frequencies, and will be used to validate model based retrieval algorithms.

1. Introduction

Accurate knowledge of snow properties is essential for snow hydrology. Onset of snow melt, snow covered area, snow wetness, and snow water equivalent are variables that goes into the hydrological models to predict run-off. These predictions are used for flood forecasting and production planning in hydro power plants. Currently 98% of the Norwegian electricity production comes from hydro power and approximately 50% of this comes from melted snow. With the establishment of a common Nordic power trading market in the late nineties, the power producers, power traders and the regulatory branches have developed a need for accurate prediction of run off. This has created a market for accurately derived snow products based on satellite data. The most important parameter is the distribution of snow water equivalent (SWE). Most of the power companies and trading companies in Norway run their own HVB (Bergström, 1992) hydrological models and for the satellite products to be useful the accumulated accuracy within a catchment must be 10% or better. One will also need to know the altitude distribution of snow within the catchment. The most important period is from onset of snow melt in the spring to the end of melt (April-June in Norways mountainous regions).

2. Current Status of Operational Snow Monitoring

Currently the only fully operational satellite based snow mapping in Norway is snow covered area (SCA) maps based on optical data (AVHRR). To improve the quality and both the spatial and temporal resolutions a combined method of SCA retrieval based on both SAR and optical (ASAR and MODIS) data has been developed (Solberg et al., 2004). These methods have been incorporated into a semi-automated production line for operational use and were tested operationally at Kongsberg Satellite Services in spring 2005. The main
reason for using SAR data for SCA mapping is the problem of lack of cloud free optical data. Clouds can be a very persistent problem for optical remote sensing in Norway.

For SAR data to be useful in an operational snow monitoring, the wideswath/scansar mode data has to be used to get large and frequent enough coverage. This will give a full coverage 2-3 times per week over Scandinavia.

2.1. SCA Retrieval Technique

The retrieval of SCA from SAR is based on the difference in backscatter between a reference scene in the same orbital track as the scene that will be classified, obtained during cold, dry snow conditions. We then classify wet snow based on the difference in backscatter caused by the high absorptance of the wet snow pack. A threshold of -3dB has been chosen based on comparison with field data (Nagler et al., 2000; Storvold et al., 2005). In the right panel of Figure 1 each classified pixel is assigned a quality value ranging from 0-100 depending on distance from threshold, viewing geometry, and quality of reference image (probability that the snow in the reference image is dry).

Combining the SAR data with temperature fields allows us to estimate quality of the retrieved data and make an assessment of the distribution of dry snow that is not directly detectable by SAR. The temperature fields are constructed based on more than 300 meteorological stations scattered throughout the region.

Dry snow is postulated in areas of higher elevation than the mean wet snow elevation within 20 km of the classified pixel at the same time as the air temperature is below freezing. The dry snow classification is assigned a quality number ranging from 0-70 depending on the elevation and air temperature. Figure 1 shows the differential backscatter between a cold winter reference scene and the May 5th 2004 scene covering most of the mountainous regions of South Norway (left panel), the resulting SCA map with wet and dry snow (center panel), and the corresponding confidence values (right panel).

2.2. SCA Retrieval Results 2004

The SAR wideswath scenes are georeferenced within a fraction of a pixel accuracy with Norut IT’s automated geocoding routine (Lauknes et al., 2005). The difference between South Norway SCA classifications based on SAR and optical data (MODIS), for the 2004 season, is shown in Table ???. Only cloud free pixels are compared and this causes the apparent large day to day variation in total SCA shown in the table. We see from Table ?? that the total snow covered area is on average approximately 4-5% larger for the Modis retrievals than for the SAR retrievals. The largest difference coincided with a late May cold period with approximately 5 cm of new dry snow.

2.3. Current Pitfalls in SAR SCA Retrievals

The main uncertainties in today’s SAR based snow covered area algorithms lies in the dry snow cover estimate,
which currently are based on air temperature and presence of wet snow from which a lower altitude boundary for the dry snow coverage is derived. This causes occasional problems in early spring when there are no wet snow present within the preset distance of the pixel that is to be classified, causing the algorithm to predict bare ground. The preset distance is chosen based on the climatic scales of the region. In the late season scattered wet snow pixels can cause false classifications of dry snow pixels during short cold spells.

The algorithm currently in use does not work in forested areas. Algorithms for estimation of snow cover in forested areas has been demonstrated in Finland (Koskinen et al., 1999) but these cannot easily be used in Norway due to the different climate regime (Norway has a coastal climate with multiple freeze-thaw cycles throughout the winter).

Compared with SCA retrieval from optical data, the SAR derived algorithm is binary (snow/no snow) and tend to underestimate the snowcover when pixels are partially snow covered. This is due to the high sensitivity to strong scatters (rocks and vegetation) that is often exposed early in the melt.

<table>
<thead>
<tr>
<th>Date</th>
<th>Snow Cover [%]</th>
<th>Date</th>
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<th>Difference [%]</th>
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3. Snow Water Equivalent

3.1. Repeat Pass Interferometry

Use of repeat pass interferometry can be used to detect changes in SWE between successive passes (Guneriusen et al., 2001). Main limitation on this method currently is the problem of unwrapping the phase change when the change in snow water equivalent is larger than typically 1-2 wavelengths. The equation below relates the phase change with the change in SWE (Guneriusen et al., 2001)

\[
\Delta \Phi_s \approx \frac{1.6}{\cos \theta_i} k \rho \Delta Z_s = \frac{1.6}{\cos \theta_i} k \Delta SWE ,
\]

where \( SWE = \rho Z_s \). For currently available SAR data (C-band) and repeat cycles (35 days Envisat) this is not a feasible approach, but with the launch of ALOS L-Band SAR this method is likely to become more useful. Still there are problems with large snowfalls and with high absorption if the snow is wet. Wet snow absorption limits the usefulness of this method in the most interesting period, which is during the snow melt.

3.2. \( \Delta k \) Repeat Pass Interferometry

For a snow density of 0.3 kg/dm\(^3\), radar wavelength of 5.62 cm (Envisat), a nd incidence angle \( \theta_i = 23^\circ \), phase wrapping occurs at a snow depth of only 10.7 cm. Retrieval of SWE from C-band SAR can be performed using repeat pass interferometry and delta-\( k \) processing. This method was demonstrated using ERS data (Engen et al., 2004) and Envisat ASAR data (Larsen et al., 2005). Based on the delta-k principle known from the radar literature (T. Hagfors, 1961), it has been proposed to handle the phase unwrapping problem by splitting the bands of both the summer and the winter image into two subbands in the slant range dimension. This results
in two bands with slightly different carrier frequencies. By forming interferograms for each of the subbands, we get two interferograms with a different phase, due to the different carrier frequencies. This phase difference, the delta-k interferometric phase, is given by

$$\Delta \Phi_s = \frac{1.6}{\cos \theta_i} (k_2 - k_1) \Delta SWE = \frac{1.6}{\cos \theta_i} \Delta k \Delta SWE,$$

(2)

where $\Delta k = k_2 - k_1$ is the difference in wavenumbers between the two subbands. For Envisat ASAR data with a maximum bandwidth of 15 MHz, we split the band into two subbands with a carrier frequency difference of about 6.5 MHz. This results in phase wrapping at a snowdepth of about 85 m. Thus, the phase unwrapping problem is effectively avoided. The largest drawback of this method is the course resolution of the SWE product due to the need for averaging to improve the signal to noise (5 x 5 km for Envisat). In Figure 2 we show a comparison between field measurements of SWE and SWE derived from Envisat.

Figure 2: Interpolated SWE field data versus ASAR SWE result. Outliers are eliminated.

3.3. Snow Volume Backscatter

Ku-band SAR instruments have the potential to overcome some of the problems that L-X-band radars have in retrieving SWE. The main advantage of Ku-Band is the large sensitivity to dry snow. Ku-band signal has sufficient penetration in most snow packs, at the same time as the volume scattering signal is detectable for dry snow. Higher frequencies have shorter penetration depths, lower frequencies have less sensitivity. According to Shi et al., [2003] Ku-band is thus the optimal sensor frequency for estimation of SWE due to the balance of detectability vs. penetration depth. Repeat pass corrections will allow for removal of topographic and terrain effects and polarimetric measurements will enable decoupling of the snow volume scattering and snow surface scattering contributions.

4. Future Snow Monitoring

There are several new upcoming SAR missions (see Table 2) over the next couple of years that yield new possibilities in snow property retrieval and have potential in operational snow monitoring. The introduction of polarimetric sensors as well as new frequencies opens new possibilities in particular in regard to operational retrieval of SWE.

4.1. Users and Requirements

Snow properties are particularly important for hydropower producers, power traders and regulators. Both for optimizing hydropower production and for avoiding flooding. For the data to be useful for this group the quality of the products has to be good (within 10% on a basin wide scale) and the coverage frequent, in particular in melting season (at least weekly).
Table 2: Future and current SAR sensors for research and operational use.

<table>
<thead>
<tr>
<th>Sensor</th>
<th>Satellite</th>
<th>f(GHz)/Polarization</th>
<th>Resolution [m]</th>
<th>Swath [km]</th>
</tr>
</thead>
<tbody>
<tr>
<td>SAR</td>
<td>Radarsat1(1995-)</td>
<td>5.3 VV</td>
<td>10, 30, 100</td>
<td>100-500</td>
</tr>
<tr>
<td>ASAR</td>
<td>Envisat (2002-)</td>
<td>5.3 HH,VV,HV</td>
<td>30, 100</td>
<td>100-400</td>
</tr>
<tr>
<td>PALSAR</td>
<td>ALOS (2006-)</td>
<td>1.2 PP</td>
<td>15/100</td>
<td>40-350</td>
</tr>
<tr>
<td>TerraSAR-X</td>
<td>TerraSAR(2006-)</td>
<td>9.6 PP</td>
<td>1, 3, 16</td>
<td>5, 30, 100</td>
</tr>
<tr>
<td>SAR</td>
<td>Radarsat2(2006-)</td>
<td>5.3 PP</td>
<td>3, 10, 25, 50, 100</td>
<td>20-500</td>
</tr>
<tr>
<td>C-SAR</td>
<td>3-Constellation</td>
<td></td>
<td>5.3 VV</td>
<td>50</td>
</tr>
<tr>
<td>Sentinel-1</td>
<td>GMES, ESA-EC</td>
<td>5.3 PP ?</td>
<td>25-50 ?</td>
<td>350-500 ?</td>
</tr>
</tbody>
</table>

PP - Polarimetric; Source: H. Rott, personal correspondance

5. Conclusion

Future sensors and new retrieval methods will allow for establishment of operational monitoring of SCA and SWE, for hydrological purposes based on SAR, that will meet the user requirements on quality and coverage. Main obstacle today is lack of operational sensors, in particular for the retrieval of SWE. Scheduled missions within the next couple of years will change this. Wet snow remains a challenge preventing melt season estimates of SWE.

Acknowledgment

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REFERENCES